Application of generalised inverse for analysis of magnetic anomalies due to a dyke Model -

- Some numerical experiments

Technical Report No: NIO/TR/07/2005

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ABSTRACT

Forward and Inverse methods of magnetic interpretation have been used to estimate the parameters of dyke like intrusion over the eastern continental shelf of India. In the forward problem, the tentative parameters of the causative source are assumed and anomalies are calculated using the anomaly equation, while inverse problem refers to methods of tracing the boundaries or outlines of anomalous bodies using an iterative approach. While solving the inverse problem, data kernel has been generated through the model (i.e. partial derivatives of magnetic anomaly function with respect to model parameters at each station forming the kernel). In solving the coupled eigenvalue problem of covariance matrices of data kernel, Singular Value Decomposition (SVD) has been performed to build Generalised Inverse Operator (GIO). This GIO is operated on the observed anomaly (with reference to the calculated) to yield improved model parameters. Data and model resolution matrices are computed to check the correctness of the solution and further analysis. The marine magnetic total field anomaly of continental shelf of Visakhapatnam is interpreted using Generalised Inverse (GI) technique, which revealed the causative source of the anomaly as a dyke model at a depth of 120 m below seabed.

The software routine details are given in Annexure-I and sample input/output and FORTRAN coding are given in Annexure-II.

Key words: Magnetic anomaly, forward problem , inverse problems, GIO, SVD, data resolution, model resolution, closeness ratio, dyke intrusive, Visakhapatnam.

Introduction

Among the many geophysical techniques, magnetic method is easy to execute and economically viable and therefore widely used in geophysical exploration. A general procedure is to assume most plausible shape for the magnetic anomaly source and compute its magnetic effect at the surface and to modify the model progressively until a reasonable fit with the observed anomaly is obtained ⁽¹⁻³⁾. The refined model represents a possible solution. The process of making inferences about the causative source from observed data is what is widely known as the inverse problem. Backus & Gilbert⁴ made an attempt for exploring the mathematical structure of the inverse problems.

The marine magnetic data that were collected in the Bay of Bengal, are interpreted in terms of the geological structures, causing the magnetic anomalies ^{5,6}. In the present work, marine magnetic total field anomalies off Visakhapatnam, east coast of India (Fig. 1) inverted by the generalised inverse approach, to yield the intrusive dyke model parameters.

Ultimately the above physical problem gives a mathematical relationship between the theoretical framework (*M*: kernel) to observed data (*dr*) and leading to system of equations (i.e., M dy = dr) for solving parameters ($dy = M_a^{-1} dr$, where M_a^{-1} is Generalised Inverse Operator (GIO)). In general the exploration problems are two types: 1) over-determined and 2) under-determined. The trivial solution of even-determined case is of rare occurrence. In all the cases a good resolution of both the model and data are desirable. The least square method solves the overdetermined problem minimizing the squared Euclidean length of the error and has a perfect model resolution. Instead, the under-determined problem can be solved with a perfect data resolution by minimizing the model length. In nature, the problems are neither purely over-determined nor purely under-determined, they are mixed type. The entire field or observation over which the determination of the parameters of the model is to be carried out comes up in observation blocks, some of which are over-determined while some are under-determined leading to non-uniqueness of the solution. The generalized inverse technique solves such a mixed determined problem over the whole region of observations and has both data and model resolutions intermediate between the two extremes ⁷. Generalised Inverse (GI) technique through Singular Value Decomposition (SVD) is employed to solve such a mixed determined problem ⁷. The advantages of involving SVD in construction of the GIO are viz., (1) SVD is objective and does not impose a predetermined form to the data; (2) It provides an objective means of ranking uncorrelated modes of variability to determine weak signals or noise from the data; and (3) it provides the modes of variability which are not correlated with one another ^{7,8}.

Earlier, Malleswara Rao¹ *et,al.*, suggested interpretation of magnetic anomalies due to bodies of polygonal cross section using Marquardt⁹ optimisation technique. In the present study, magnetic anomalies due to intrusive dyke model are analysed using the approach of Natural Generalised Inverse (GI) technique ¹⁰ as it throws light on data and model resolution and individual model parameter contribution. All these procedures and algorithms attempt to relate physical models to readily observable data and determine the "best-estimates" of model parameters.

Notation

- *J* intensity of effective magnetization of two-dimensional body lying in a plane perpendicular to the strike
- ϕ dip of the effective magnetization
- $lpha\,$ strike of the body measured from the magnetic north due east or west
- *i* dip of the earth's magnetic field
- F Earth's magnetic field expressed in gammas
- D_m direction of measurement, is 0 for horizontal ΔH , $\Pi/2$ for vertical ΔV and i for the total field ΔF anomalies.
- ΔT anomaly in any component : it is equal to ΔH , ΔV and ΔF when D_m is 0, $\Pi/2$ and i respectively.

 ΔT_{cal} - calculated anomaly

- dT ($\Delta T_{obs} \Delta T_{cal}$)
- *K* susceptibility contrast of the body
- θ geological dip of the dyke model
- Q an angle, which is a function of $\theta,\,\phi,\textbf{D}_{m},\,i$ and α
- distance from an arbitrary reference point to the position of centre of the dyke
- *x* distance co-ordinates of anomalies
- w half width of the dyke
- *C* size factor, being a function of magnetisation(in gammas) and dip of dyke
- *z* depth to the top of the dyke
- N_{OBS} Number observations
- A' regional gradient
- B' regional at the first station

Magnetic dyke - forward problem

Dykes and thin sheets are frequently used geophysical models in magnetic studies (Radhakrishna Murthy¹¹⁻¹³. The dyke is assumed to be a model with a horizontal top bounded by parallel sides (Fig. 2). Equation for its magnetic anomaly can be written by treating it as a model bounded by four first order surfaces. The conventional method is to write down the magnetic anomaly of a line dipole of cross-sectional area du . dv and integrating it throughout the cross-section of the dyke. When the origin is placed above the center of the dyke the generalized equation for its magnetic anomaly in any component can be written as (Radhakrishna Murthy¹²):

$$\Delta T(x) = 2J' \int_{z_1 - w + (v - z_1) \cot \theta}^{z_2} \int_{(v - z_1) \cot \theta}^{w + (v - z_1) \cot \theta} \frac{(v^2 - (x - u)^2) \sin \varphi' - 2(x - u)v \cos \varphi' dv du}{((x - u)^2 + v^2)^2}$$
------(1)

where *w* is the half-width of the dyke, θ is its dip, z_1 and z_2 are the depths to its top and bottom and

$$J' = J(1 - \cos^2 \alpha \ \cos^2 D_m)^{1/2},$$

$$\phi' = \phi - \arctan(\sin \alpha \ \cot \ D_m)$$

with *J*, ϕ , α and *D_m* as defined in notation. The usefulness of the above equation is that one can use the same for calculating vertical, horizontal or total field anomalies by suitably taking the values of *D_m*. Equation 1 after integration simplifies to

$$\Delta T(x) = 2 J' \sin \theta \left[\cos Q \cdot (\theta_A - \theta_B + \theta_C - \theta_D) + \sin Q \ln (r_B r_D / r_A r_C) \right] -----(2)$$

where various parameters are defined in Fig.2. For special case when the dyke is extending to a large depth $\theta_C = \theta_D$ and $r_D = r_{C.}$ Thus, Eq.(2) takes the form (Radhakrishna Murthy³):

The distances and angles involved in Eq. (3) are defined as follows.

$$\theta_A - \theta_B = \arctan((x+w)/z) - \arctan((x-w)/z)$$
, for $z \neq 0$,

$$r_A^2 = (x+w)^2 + z^2$$
, $r_B^2 = (x-w)^2 + z^2$

with $Q = \theta - \phi$ and $C=2 J' \sin \theta$.

The inverse problem

The generalised equation for the magnetic anomaly of a dyke model is of infinite depth extent shown in Eq:(3). The parameters to be determined from its anomalies are *w*, *z*, *Q* (from θ - ϕ), *J* sin θ (from *J*' sin θ) and *D* the position of the center of the dyke. For this, the anomalies $\Delta T(x_i)$ are digitized at distances measured from a convenient reference on the profile. When the distances are measured relative to the reference point, the anomaly equation is rewritten as

 $\Delta T(x_i) = C [\cos Q (\theta_A - \theta_B) + \sin Q \ln (r_B / r_A)] + A'x_i + B', \qquad ------(4)$

where $A'x_i + B'$ is the associated regional anomaly. Initialization of these parameters is achieved by the procedure suggested by Radhakrishna Murthy¹⁴, which considers the dominant maximum and minimum anomalies and their corresponding distances from the chosen reference point on the magnetic anomaly profile. To calculate the initial values these characteristic distances and/ or anomalies are scaled by the computer from the field profile and their ratios or magnitudes are used to calculate the parameters by properly derived empirical relations. The characteristic distances to be measured are chosen, by trial and error, such that they or their ratios vary significantly with one of the parameters and less significantly with others.

Most often the anomalies ΔT_{cal} computed involving these initial parameters differ from the observed anomalies ΔT_{NOBS} . The difference between observed and calculated anomalies at the *i*th anomaly point is expressed by the Taylor's expansion in terms of errors in the initial values of the model parameters which can be expressed, in terms of errors *dw*, *dz*, *dD*, *dQ*, *dC*, regional gradient (*A'*) and constant (*B'*), in the parameters as follows:

$$dT(x_i) = \Delta T_{obs}(x_i) - \Delta T_{cal}(x_i) \qquad (i=1, \dots N_{OBS}) \quad ------(5)$$

$$7$$

$$dT(x_i) = \Sigma \quad \underline{\partial \Delta T}(x_i) \, da_K,$$

where $da_1 = dw$, $da_2 = dz$, $da_3 = dD$, $da_4 = dQ$, $da_5 = dC$, $da_6 = dA'$, and $da_7 = dB'$.

The partial derivatives are calculated through the following equations: (Radhakrishna Murthy¹³):

$$\begin{aligned} \frac{\partial \Delta T(x_i)}{\partial a_1} &= 2C \Biggl[\frac{\left[(x_i - D)^2 + z^2 + w^2 \right] z \cos Q - (x_i - D) \left[(x_i - D)^2 + z^2 - w \right]^2 \sin Q}{\left[(x_i - D + w)^2 + z^2 \right] \left[(x_i - D - w)^2 + z^2 \right]} \Biggr], \\ \frac{\partial \Delta T(x_i)}{\partial a_2} &= 2Cw \Biggl[\frac{\left[(x_i - D)^2 + z^2 - w^2 \right] \cos Q + 2z(x_i - D) \sin Q}{\left[(x_i - D + w)^2 + z^2 \right] \left[(x_i - D - w)^2 + z^2 \right]} \Biggr], \\ \frac{\partial \Delta T(x_i)}{\partial a_2} &= 2Cw \Biggl[\frac{2z(x_i - D) \cos Q - \left[(x_i - D)^2 - z^2 - w^2 \right] \sin Q}{\left[(x_i - D + w)^2 + z^2 \right] \left[(x_i - D - w)^2 + z^2 \right]} \Biggr], \\ \frac{\partial \Delta T(x_i)}{\partial a_4} &= C \Biggl[\int \arctan\left((x_i - D - w)/z \right) - \arctan\left((x_i - D + w)/z \right) \int \sin Q \\ + 0.5 \cos Q \ln \left[\left((x_i - D - w)^2 + z^2 \right) / \left((x_i - D + w)^2 + z^2 \right) \right] \Biggr], \\ \frac{\partial \Delta T(x_i)}{\partial a_5} &= C \Biggl[\int \arctan\left((x_i - D + w)/z \right) - \arctan\left((x_i - D - w)/z \right) \int \cos Q \\ + 0.5 \sin Q \ln \left[\left((x_i - D - w)^2 + z^2 \right) / \left((x_i - D + w)^2 + z^2 \right) \right] \Biggr], \end{aligned}$$

 $\frac{\partial \Delta T(x_i)}{\partial a_6} = x_i ,$

and

$$\frac{\partial \Delta T(x_i)}{\partial a_i} = 1.0 \quad . \tag{6}$$

The necessary normal equations to be solved for the increments da_{K} (k=1, ...7) are framed involving Eq.(3), the observed data and partial derivatives with respect to model parameters and are given by

$$\sum_{i=1}^{N_{OBS}} \sum_{k=1}^{7} \frac{\partial \Delta T(x_{i})}{\partial a_{j}} da_{k} \frac{\partial \Delta T(x_{i})}{\partial a_{k}} (1 + \delta \lambda) da_{k}$$

$$= \sum_{i=1}^{N_{OBS}} \frac{\partial \Delta T(x_{i})}{\partial a_{j}} dT(x_{i}) , \qquad (j = 1 \dots 7)$$

with N_{OBS} = number of magnetic anomaly points , the constant $\delta = 1$ for i = j,

and 0 otherwise and λ is Marquardt⁹ damping factor. The Marquardt damping factor λ is chosen by the relation $\lambda = 0.5 [2^{n-1} - 1]$ (Radhakrishna Murthy ¹³). The value of λ is initially set to zero by putting n=1. Subsequently, it is increased or decreased, each time, by one depending upon the resulting objective function is increased or decreased. The Eq.(7) may be be put in the matrix form:

 $M \, dy = dr \tag{8}$

The Eq. set (8) (for $i = 1, ..., N_{OBS}$) is to be solved for the increments to the various parameters. To handle this problem there are different approaches that include - a) ridge regression, frequently called Marquardt⁹ iteration and b) Generalized Inversion (GI) approach. In the present magnetic data inversion, we follow the second method viz. GI, as it has additional advantages over the Marquardt method (as already mentioned in the beginning).

Equation(8) is written in following discrete form for computational convenience:

$$dr = M dy$$
; $dr = (dr_i)$; $dy = (dy_k)$; $(i = 1, 2, ..., N_{OBS}; k = 1, 2, ..., 7)$ ------(9)

Equation (9) is solved by generalised inverse^(8,15-17) (GI) through singular value decomposition (SVD) to obtain the following optimised model parameters^{10,18} dy_P as follows:

$$M = U / V^{T}$$
(N_{OBS} x m) (N_{OBS} x r) (r x r) (r x m) (10)

The columns of U and V are orthonormal i.e $U^{T}U = I_{NOBS}$ and $V^{T}V = I_{m}$

with I_{NOBS} and I_m are unit matrices of order ($N_{OBS} \times N_{OBS}$) and ($m \times m$) respectively and m = number of model parameters.

In general, $UU^T \neq I_{NOBS}$ and $VV^T \neq I_m$; $p \leq r = min (N_{OBS}, m)$ ------ (11) where r is the rank of the matrix *M* (data kernel), p number of factors to be considered and the matrices U and V are the respective coupled Eigen vector matrices for the Eigen value problem defined as:

$$(M M^{T}) U \equiv l^{2} U$$
 -----(12)

and

$$(M^T M) V \equiv l^2 V \quad -----(13)$$

In Eq.(10), / is diagonal matrix of non-zero singular values ($l_i^2 = \lambda_i$) of *M* arranged in the decreasing order. Once *U*, / and *V* are obtained by solving above Eigenvalue problem (Eqs.12 and 13) the generalised inverse solution (Jackson¹⁸) is given by

$$dy_P = V V^T dy = M_g^1 dr$$
 -----(14)

where $M_g^1 = V_p \int_p^{-1} U_p^T$ is the natural generalised inverse operator(GIO). Let $s(d) = \{u_1, u_2, \dots, u_n\}$ and $s(m) = \{v_1, v_2, \dots, v_m\}$ be sets of orthogonal vectors (i.e., column vectors of complete data space *U* and model space *V*) as detailed below:

 $U = \{U_p, U_{r-p}, U_0\}$ is the complete data space where

- U_p : optimised data space set of p eigenvectors corresponding to dominant eigenvalues such that $\lambda'_1 > \lambda'_2 > \lambda'_3 > \dots > \lambda'_p \neq 0$.
- U_{r-p} : set of (*r*-*p*) eigenvectors corresponding to remaining dominant eigenvalues such that $\lambda'_{p+1} > \lambda'_{p+2} > \dots > \lambda'_r \neq 0$ and
- U_0 : Null space having eigenvectors corresponding to r- min ($N_{OBS} \times m$) zero eigenvalues.

Similarly it follows in the case of complete model space (i.e. $V=\{V_p, V_{r-p}, V_0\}$) and can also be explained .

If VV^{T} equals to I_{m} (i.e. the rank of the matrix M = m) then the solution of the Eq. (8) is

 $dy_P = dy = M^{-1}_{g} dr$ -----(15)

If VV^{T} not equals to I_m (for the case of, presence of noise in the model space)

$$dy_P \approx dy = M^{-1}_g dr \qquad -----(16)$$

For better estimates the resolution in the model space VV^{T} of Eq.(16) has to be improved. This is done judiciously by selecting the $p(\leq r)$ eigenvectors (also called factors or modes) in the activated model space s(m) corresponding to the ranking of the p singular values (λ'_{i}) of the data kernel (*M*) in descending order. The noise in the data kernel prevailing in the form of small values increases the ranking of the matrix apart from amplifying the solution. This however does not provide any additional information on the model parameters. So, it can be treated as though the solution to the present problem is obtained through the optimisation.

The above process is continued to obtain optimised model based on the best fit between the observed and calculated anomalies. Once the optimised solution is obtained, it is necessary to assess how best the data determines the model parameters. This is done through model resolution $(V_p \ V_p^T)$ and data resolution $(U_p \ U_p^T)$ matrices.

The following mathematical presentation shows us how the data and model space are interrelated. Pre-multiplying (Eq. 14) by V_P^T on both sides of the equation and by defining the general parameters

and

 $\alpha = V_P^T dy_P \qquad ------(17)$ $\Rightarrow \qquad \alpha_j = v_j^T dy_P$ $\beta = U_P^T dr_P \qquad -----(18)$ $\Rightarrow \qquad \beta_j = u_j^T dr_P$

it reduces to

 $\alpha = \Gamma^{-1} \beta \qquad \qquad (19)$

The relation shows how data errors in β component belong to small eigenvalues λ'_i (of diagonal matrix Γ) are strongly amplified giving rise to large errors of α_j and hence those components dy, which contributes to α_j . For the investigation of parameter errors and information content of the data it has become tradition to use the model resolution matrix $R = V V^T$ and information density matrix $S = U U^T$.

From Eq. (17) we see that an investigation of the vector v_i will reveal which

parameters of dy are specifically attached to the generalised parameters a_j . From Eq. (18) we see that similar investigation of the vector u_j will reveal which data of dr are particularly contributing to the content of the generalised datum β_j . The components of v_j directly indicate which of the parameters of dy are determined. On other hand the components of u_j indicate where to find those data points that contribute to the determination of these parameters v_i^{10} .

Closeness ratio: The ratio between the sum of the factors considered and that of the data matrix (kernel) is the measure of closeness of the model data

Measure of closeness =
$$\sum_{i=1}^{p} \lambda'_i / \sum_{i=1}^{r} \lambda'_i$$
 ------ (20)

where r is the rank of the data matrix M (Kernel). The first eigen function, u_1 associated with the largest eigenvalues λ'_1 represents the gross features (i.e., $|| u_1^T M||$) in the data in the least square sense, while the second function u_2 associated with second largest eigenvalue λ'_2 describes the residual mean square data [i.e., $|| u_2^T (M - (u_1 u_1^T M))||$] in the least square sense and so on. The closeness ratio is expressed in percentage to judge the contribution of different parameters.

SOFTWARE DETAILS:

The software program details for the dyke model analysis using GI are presented in program *GIDYKE*.FOR (**ANNEXURE – I**) with another synthetic dyke model as input and output details with FORTRAN coding in **ANNEXURE – II**.

Dyke inversion - Synthetic and field examples:

The performance of the natural Generalised Inverse (GI) explained above is illustrates with synthetic and field data sets by carrying two numerical experiments as detailed below.

Synthetic data:

A synthetic dyke model with parameters w = 4.0 km, z = 1.0 km, D = 10.0 km, $Q = 90^{\circ}$ and J = 100 are taken and computed the magnetic anomaly at 21 points (NOBS) using the forward model Eq.(3). The synthetic anomaly so generated is subjected to GI and the derived model anomaly denoted by stars and corresponding model (dotted circles) shown in Fig. 3A, which is well coincided with actual synthetic model anomaly denoted solid line. The GI optimized model parameters are w = 3.99 km, z = 1.07 km, D = 10.0 km, $Q = 90^{\circ}$ and J = 101.06 shows the efficacy of the method. Root Mean square Error (RMS) at initial (zero) iteration is 9.83 and the objective function is 2806. At 5th iteration, RMS error in GI optimized model is 1.68 while in Marquardt Inversion (MI) after 24 iterations the RMS error is 4.72 and objective function is 468 (Table I b). The success of synthetic model inversion with GI prompted to carryout the field data inversion using GI.

Field data:

Marine magnetic total field data was collected, using Baringer Proton precession magnetometer (model M123) off Lawson's Bay (opposite Kailash Hill) of Visakhapatnam(17° 42' N , 83° 17' E) along a coast parallel profile (HH') (Fig. 1) at 40 m water depth, covering a distance of 2.5 km, at one-minute polarisation

interval. Figure 3B shows the observed magnetic anomaly along this coast parallel profile HH' of Fig.1. The observed total field marine magnetic data were corrected using 1985 IGRF coefficients (DGRF¹⁹) and the residual data at 26 points on the profile is considered in the present GI analysis. The residual marine magnetic anomaly ranges between -460 to 170 nT with a low of -460nT at 1.25 km from the reference point, flanked by highs on either side, 60nT(left) and 170 nT (right)(Fig.3B). This magnetic anomaly is subjected to GI to yield optimized dyke parameters. Figure 3B also shows the calculated anomalies of initial and final dyke model using GI.

Field example - Results and discussion

In the present section, computed results of the field data (Fig.3B) are discussed. SVD of kernel $M_{(26\times7)} = U/V^T$ has been performed on residual magnetic data and shown as observed data in Fig. 3B. The column matrices $U = (u_1, u_2, ..., u_7)$ and $V = (v_1, v_2, ..., v_7)$ are orthonormal eigenvectors spanning data space s(d) and model s(m) corresponding to eigenvalues λ_i of diagonal matrix /i (Table 1) computed by solving coupled eigen value problem of two covariance matrices MM^T and M^TM of 7 components contributing 100% of total variance. Based on the point of inflection or minimum point of cumulative eigenvalues, four factors (p=4) are retained for interpretation as shown in Figs.4 and 5 (however other figures in respect of other 3 factors are also shown for the sake of completeness). The GIO $M^1_g = V_p / p^{-1} U_p^T$ has been built after removing null spaces $\{U_0, V_0\}$ to operate on observed magnetic anomaly to yield dyke intrusive model parameters in the present studies (Fig.3B, field model).

Using vectors in activated space, data resolution matrix $\{U_p \ U_p^T\}$ and model resolution matrix $\{V_p \ V_p^T\}$ have been computed and the elements of these matrices are presented in 3-D contour form (Figs.6 and 7) respectively. The diagonal elements of these resolution matrices are presented in Tables 2 and 3. The column Vectors u_i (i=1, ..., 7) of s(d) of corresponding eigenvalues of λ_i^{\prime} (i.e. measure of variance *M* in descending order), having 26 components and with each component a value that indicates amplitude, which explains sharing contribution of its variance mode *i*. The component u_i represent the trend or

direction of spatial common features contained in kernel M (partial derivatives of magnetic function with respect to model parameters). Its corresponding eigenvalue λ_i represents energy level. In a similar manner, the components of vectors in model space s(m) can be explained. The percentage contribution of individual components of s(m) and s(d) have been computed through closeness ratio approach (Eq. 20) and are shown in Tables 4 and 5. Eigenvectors derived from the data kernel of present field example have been presented in Figs. 4 and 5 to explore the spatial and model parameter variability. From these Figs. 4 and 5 and Tables 4 and 5 the first four eigenfunctions of s(d) and s(m) account for more than 98% of variance relatively dominated by the first function alone. The first function u_1 and v_1 contribute 58% of total distribution representing model parameters Q followed by D, w, and z in that order. In general, Q - an angle, function of θ , ϕ , D_m , *i* and α , contributes to the extent of 45% between stations 18 and 21. Further z < w, indicating that it is thick dyke intrusive. The contribution of remaining three functions in s(d) and s(m) is only 2% of total information, and is too insignificant to consider for interpretation.

Therefore the first four energetic spatial and model functions of s(d) and s(m) have been used to bring the predictable rich part of the original signal to obtain inverse solution (Fig.3B) by operating GI on magnetic anomaly and for construction of data and model resolution matrices in 3-D form (Figs. 6 and 7). At the end of 21 iterations the results obtained from GI indicate that the width of the dyke intrusive model is 0.72 km at a depth (*Z*) of 0.12 km from the surface located at (*D*) 1.44 km away from the reference point and at an angle (Q) of 140° and with intensity of magnetization (*C*) 225.86 nT (Fig 3B). The optimized value of intensity of magnetization (*C*) is reasonably comparable with local value (250. nT) ¹². Computed model anomaly with refined parameter set almost coincides with observed field anomaly to quantify (Fig. 3B). The initial objective function (=111039) is reduced up to (=4645) after 21 iterations, and the corresponding refined model parameter set is considered for interpretation. The refined intrusive dyke model corresponding to minimum objective function is shown in Fig.3B. The Root Mean Square (RMS) error for the field data at zero iteration is 65.35. After 21

iterations it has reduced to 13.36.

Data resolution:

The data resolution matrix is an indication of the information density of the data kernel i.e. it indicates which data contribute independent information to the solution. The diagonal elements of $U_{p} U_{p}^{T}$ are shown in Table 3 for factors 1 to 7. A value of unity indicates contribution of information independent of other observations. From Table 3 one can infer which of the data points contribute strong / poor information resolution. The data resolution is given by $N = MM^{1}_{g} = U_{p}U_{p}^{T}$. The data are completely resolved if U_{p} spans the complete space of data. The small eigenvalues in the data kernel increases the rank of the matrix (i.e. the dimension of activated space) besides amplifying the solution due to the presence of noise. Such a filtered data kernel has been utilized in 3-D form (Fig. 8), which almost represents the original data kernel to build GIO. In Fig. 6 Factor 1 describes the contour map of data resolution matrix $u_1 u_1^T$, where u_1 is the highest spatial energetic mode of s(d). It gives 58.003% variance of the total information i.e. mostly gross features of the kernel (Table 5). High resolution is found at stations 10 and 18 indicating the sudden change at these stations likely indicating the width of the dyke intrusive (Fig. 6 and Tables 3-5)

Model resolution:

The model resolution of the generalised inverse is given by $R = M_g^1 M = V_p$ V_p^T . Here p indicates number of factors used in SVD, which is less than or equal to the rank of the matrix M. The model parameters will be perfectly resolved if V_p spans the complete space of the model parameters i.e. $V_p V_p^T = I_m$. The model resolution is perfect for the contribution of magnetic anomaly due to Q (Fig. 7, factor 1 and shown in Table 4) and one can conclude that Q is independent of other model parameters and of homogeneous nature which is also reflected in the data resolution Fig. 6, factor 1 at stations 12 and 18 (also see Table 5). Table 4 presents the individual parameter contribution in each of the seven factors. In first factor the contribution of Q is 45.222% out of 58%, and in second factor the contribution of W is 9.7% out of 18.5%, and in third factor z contributes 7.1% out of 13.2% and finally in fourth factor that of D is 5.5% out of 8.1%. The total contribution of Q, w, z and D is 62% out of 89.8% from the first four factors.

From the above model and data resolution study, the partial information contained in the data space is adequate enough to reconstruct the approximate model parameters. This reveals the efficacy of natural generalised inverse GI with a least square sense in handling problems of over determinacy.

Acknowledgements:

The authors are thankful to Director, National Institute of Oceanography, Dona paula, Goa for his keen interest and encouragement.

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Synthetic Example – Dyke Model:

Table: I (a) Analysis by generalised inverse

Magnetic anomalies - due to dyke model

ITERATION NUMBER 0 _____ OBSERVED CALCULATED ERROR ANOMALY ANOMALY DIS-TANCE _____ .00 1.00 2.00 3.00 4.00 5.00 5.50 6.00 6.50 7.00 7.50 8.00 8.50 9.00 10.00 11.00 11.50 12.00 12.50 13.00 13.50 14.00 14.50 15.00 16.00 17.00 18.00 19.00 20.00 _____

OBJECTIVE FUNCTION IS 2806.

MODEL PARAMETERS

DEPTH TO TOP OF DYKE IS	1.83
HALF THE THICKNESS IS	3.66
THE ORIGIN IS AT	10.00
THETA-PHI IS	90.00
CONST IS	136.09
REGIONAL AT 1ST STATION IS	-17.81
REGIONAL GRADIENT IS	1.78

DIS- TANCE	OBSERVED ANOMALY	CALCULATED ANOMALY	ERROR
.00	83.6	83.4	.2
1.00	93.9	93.9	.0
2.00	107.2	107.4	.2
3.00	125.1	125.4	.3
4.00	150.3	150.4	.1
5.00	185.7	184.3	1.4
5.50	197.5	200.5	3.0
6.00	208.7	204.7	4.0
6.50	185.0	188.5	3.5
7.00	160.9	159.9	1.0
7.50	130.0	129.3	.7
8.00	100.1	100.3	.2
8.50	74.0	73.4	.6
9.00	47.8	48.0	.2
10.00	.0	.0	.0
11.00	-47.8	-48.0	.2
11.50	-74.0	-73.4	.6
12.00	-100.1	-100.3	.2
12.50	-130.0	-129.3	.7
13.00	-160.9	-160.0	.9
13.50	-185.0	-188.5	3.5
14.00	-208.7	-204.7	4.0
14.50	-197.5	-200.5	3.0
15.00	-185.7	-184.3	1.4
16.00	-150.3	-150.4	.1
17.00	-125.1	-125.4	.3
18.00	-107.2	-107.4	.2
19.00	-93.9	-93.9	.0
20.00	-83.6	-83.4	.2

Table I(a)ITERATION NUMBER 5

OBJECTIVE FUNCTION IS 83.

MODEL PARAMETERS

DEPTH TO TOP OF DYKE IS	1.07
HALF THE THICKNESS IS	3.99
THE ORIGIN IS AT	10.00
THETA-PHI IS	90.00
CONST IS	101.68
REGIONAL AT 1ST STATION IS	-1.32
REGIONAL GRADIENT IS	.13

Table : I (b) GI analysis - RMS error details

Model	Z	W	D	$Q = (\theta - \phi)$	CONST.	Reg.	Itr /	RMS
							Obj.fun	err
Exact	1.0	4.0	10.0	90.0	100	0		
Initial	1.83	3.66	10.0	90.0	136.09	-17.8	0/2806	9.83
GI Mdl.	1.07	3.99	10.0	90.0	101.68	-1.32	5/83	1.69
MI Mdl.	1.0	4.01	9.99	89.7	119.2	-23.4	24/468	4.72

Field example - Dyke model:

Table 1: Eigen values and corresponding closeness

ratio for Magnetic dyke model - Field Example .

Eigen	Closeness
Values	Ratio
35.97	58.003
20.36	76.576
17.18	89.804
13.45	97.915
6.60	99.869
1.32	99.946
1.09	100.000

Table 2: Diagonal elements of model resolution ((V _p V	Ъ ^Т))
	4 P	P /	/

Model/ Factor	W	Z	D	Q	С	A	В
1	0.0003	0.0001	0.0513	0.9476	0.7807	0.0007	0.21E-6
2	0.6563	0.3340	0.0535	0.9477	0.0002	0.0083	0.00005
3	0.9824	0.9996	0.0550	0.9477	0.0002	0.0149	0.00009
4	0.9876	0.9996	0.9823	0.9968	0.0002	0.0331	0.00009
5	0.9999	0.9998	0.9999	0.9999	0.0015	0.9927	0.00590
6	1.0000	0.9999	0.9999	1.0000	0.7568	0.9976	0.24560
7	1.0000	1.0000	1.0000	1.0000	1.0000	1.0000	1.00000

Parameters	v ₁ %	V ₂ %	V ₃ %	V ₄ %	V ₅ %	v ₆ %	V ₇ %
W	.837	-9.693	4.972	427	.153	.000	.000
Z	186	6.917	7.106	046	.021	.001	.000
D	10.504	.562	340	-5.563	185	.000	.000
Q	- 45.222	108	.009	-1.282	078	.000	.000
Const	.000	168	026	.000	050	.046	019
A	-1.255	1.041	714	780	1.361	.004	.002
В	.000	.084	061	012	.107	026	033
Total (%)	58.003	18.573	13.228	8.111	1.454	0.0078	0.054

Table 4: Percentage contribution of individual components of model space s(m)

Station ID	u ₁ %	u ₂ %	u ₃ %	U4%	u ₅ %	u ₆ %	u ₇ %
0	1.004	.079	138	.233	.042	003	006
1	1.121	.085	0157	.245	.054	003	005
2	1.223	.095	182	.258	.065	002	004
3	1.325	.106	213	.270	.077	001	003
4	1.465	.116	251	.284	.088	001	002
5	1.618	.137	307	.296	.098	.000	001
6	1.821	.169	383	.299	.106	.000	001
7	2.076	.227	499	.288	.109	.000	.002
8	2.420	.349	681	.227	.098	002	.003
9	2.891	.671	966	.023	.052	003	.004
10	3.401	1.612	-1.217	526	075	.003	.001
11	3.070	2.732	242	757	104	.003	004
12	1.554	1.981	.725	117	.068	003	.001
13	.153	1.379	.809	.036	.107	006	.003
14	-1.006	1.173	.743	018	.087	006	.003
15	-2.089	1.199	.668	074	.053	003	.001
16	-3.184	1.437	.546	006	.016	.004	001
17	-4.330	1.892	.107	.415	018	.005	002
18	-4.955	1.580	-1.107	.869	041	.005	.001
19	-4.114	.132	-1.245	.000	028	006	.001
20	-3.210	275	753	456	.016	006	001
21	-2.611	301	464	536	.056	002	002
22	-2.203	264	311	523	.088	.001	001
23	-1.911	227	220	487	.114	.003	001
24	-1.694	190	166	450	.136	.005	.000
25	-1.516	164	129	417	.156	.007	.001
Total(%)	58.003	18.573	13.228	8.111	1.954	.078	.054

Table 5: Percentage contribution of individual component of data space s(d)

ITERAT	TION NUME ↓	BER 0		ITERATION NUMBER 21 \downarrow		
DIS- TANCE	OBSERVE ANOM	D CALCULATI MALY (nT) ERF	ED R.(nT)	CALCULATED ANOMALY(nT) ERR. (nT)	
0.00	60.	148.	88.	56.	4.	
1.00	55.	122.	67.	53.	2.	
2.00	52.	95.	43.	49.	3.	
3.00	49.	66.	17.	44.	5.	
4.00	40.	34.	6.	38	2.	
5.00	20.	-1.	21.	29.	9.	
6.00	10.	-41.	51.	17.	7.	
7.00	-5.	-87.	82.	-1.	4.	
8.00	-30.	-141.	111.	-30.	0.	
9.00	-70.	-205.	135.	-81.	11.	
10.00	-160.	-281.	121.	-180.	20.	
11.00	-350.	-359.	9.	-337.	13.	
12.00	-455.	-409.	46.	-436.	20.	
13.00	-460.	-410.	50.	-456.	4.	
14.00	-450.	-375.	75.	-444.	6.	
15.00	-407.	-323.	84.	-412.	5.	
16.00	-335.	-261.	74.	-358.	23.	
17.00	-245.	-190.	55.	-261.	16.	
18.00	-108.	-109.	1.	-95.	13.	
19.00	35.	-17.	52.	46.	11.	
20.00	115.	77.	38.	102.	13.	
21.00	100.	145.	45.	124.	24.	
22.00	110.	166.	56.	133.	23.	
23.00	130.	153.	23.	137.	7.	
24.00	150.	128.	22.	140.	10.	
		At	Iterat	ion 0	At Iteration 21	
OBJEC	TIVE FUN	ICTION IS	11	1039.	4645.	
MODE		ETERS				
DEPTH	I TO TOP	OF DYKE IS	2	2.25 units	0.12 units	
HALF		KNESS IS	4	.50 units	0.36 units	
THE O	RIGIN IS A	AT	16	.10 units	1.44 units	
IHEIA	-PHI IS		86	5.05°	140.78°	
CONS	I IS		67	(.58 n l	225.86 nT	
REGIO	NAL AT 1	ST STN. IS	28	9.63 nT	95.15 nT	
REGIO	NAL GRA	DIENT IS	-15	5.39	1.01 nt	

Table 6: INTERPRETATION OF MAGNETIC ANOMALIES DUE TO A DYKE BY GENERALISED INVERSE APPROACH

ANNEXURE - I

SOFTWARE PROGRAM DETAILS:

Main Program GIDYKE.FOR

The main program GIDYKE.FOR in FORTRAN – 77 code form consists of seven routines GIO, SVD, MATMLT, TRNSPZ, ARRANGE, OUTPUT and the seventh DYKE.FOR (DYKE forward model). The necessary input to the main program is given in file DYKEIN.DAT that consists of number of anomaly (data) points: NOBS; number of iterations to carry out: NITER followed by corresponding model details. The difference between observed and model anomaly is computed and stored in array ERROR (K) and sum of squares of these errors is stored as single variable FUNCT (objective function). Now data kernel (coefficient matrix) is generated in matrix P of order (N_{OBS} x M) and column matrix PB of order (N_{OBS} x 1), consisting of difference between observed and computed anomaly are also stored. The output file DYKEOUT.DAT consists of model parameter details for each iteration and objective function. The main program opens files like DYMRES.DAT & DYDRES.DAT to write model and data resolution details for each iteration.

At this stage subroutine GIO is called to obtain output argument Q as column matrix of order ($N_{OBS} \times 1$) by sending input arguments P and PB. The model parameters are updated and fed to DYKE.FOR to compute forward anomaly and there by FUNCT1 which is the current objective function. In the main program this process is repeated iteratively such that ensemble average of the squares of the difference between estimated field and true field at each point is minimum, to

compute the best model parameters.

Subroutines:

Subroutine GIO: This subroutine GIO takes the data kernel A or order ($N_{OBS} \times M$) and D is a column matrix of order ($N \times 1$) consisting of deviation in the anomaly from the main program and calls the subroutine SVD to perform singular value decomposition to get orthonormal matrices U of order ($N_{OBS} \times NF$), diagonal matrix W of order ($NF \times NF$) and V matrix of order ($N_{OBS} \times NF$). It calls subroutine ARRANGE to arrange eigenvalues in descending order and corresponding vectors are arranged in GIO. Then it computes the closeness ratio as described in theory (Eq. 20) after that it generates generalised inverse operator matrix, X. Using U, W and V matrices and it operates on D matrix and computes the inverse solution and stores in column matrix DELC ($N \times 1$). Also it computes the data resolution, model resolution and returns to the calling program.

Subroutine SVD: This subroutine SVD takes data kernel A (N_{OBS} x N) from GIO and computes covariance matrix to solve the eigen value problem to perform SVD and to yield orthonormal matrices U, V and diagonal matrix W. The routine LINPACK available in public domain is used in SVD routine.

Subroutine MATMLT: This subroutine MATMLT is called from the calling program to compute the multiplication of two matrices A and B of order (N_{OBS} x N) and (N x P) and stores in Matrix C of order (N x P).

Subroutine TRNSPZ: This subroutine TRNSPZ is called in the calling program many times to compute the transpose of matrix A of order (N x N_{OBS}) and stores in matrix B of order (N_{OBS} x N).

Subroutine ARRANGE: This subroutine ARRANGE is called in subroutine GIO and arranges eigen values in descending order.

Subroutine OUTPUT: This subroutine OUTPUT is simply an output subroutine that writes the SVD results to an access disk file (SVDOUT.DAT).

ANNEXURE - II

Synthetic model

GIODYKE PROGRAM INPUT AND OUTPUT FILE DETAILS

INPUT - FILE: DYKEIN.DAT

17,45.0,90.0,25 0.0,1.0,2.0,3.0,4.0,5.0,6.0,7.0 8.0,9.0,10.0,11.0,12.0,13.0,14.0,15.0 16.0 -220.,-300.,-370.,-460.,-495.,-340.,-210.,-80. 60.,230.,390.,510.,560.,450.,330.,250. 200.

OUTPUT - FILE: DYKEOUT.DAT

ITERATION NUMBER 0

ANALYSIS BY GENERALISED INVERSE

INTERPRETATION OF MAGNETIC ANOMALIES DUE TO A DYKE USING ITERATIVE APPROACH. THE FOLLOWING TABLES PROVIDE INFORMATION ON ANOMALIES AND MODEL PARAMETERS AT THE END OF EACH ITERATION.

DIS- TANCEOBSERVED ANOMALYCALCULATED ANOMALYERROR ANOMALY.00-220180.40.1.00-300250.50.2.00-370330.40.3.00-460416.44.4.00-495469.26.5.00-340419.79.6.00-210283.73.7.00-80126.46.8.0060.31.29.9.00230.188.42.10.00390.346.44.11.00510.481.29.12.00560.528.32.13.00450.473.23.14.00330.387.57.15.00250.307.57.16.00200.238.38.				
$\begin{array}{cccccccccccccccccccccccccccccccccccc$	DIS- TANCE	OBSERVED ANOMALY	CALCULATED ANOMALY	ERROR
$\begin{array}{cccccccccccccccccccccccccccccccccccc$.00 1.00 2.00 3.00 4.00 5.00 6.00	-220. -300. -370. -460. -495. -340. -210.	-180. -250. -330. -416. -469. -419. -283.	40. 50. 40. 44. 26. 79. 73.
15.00 250. 307. 57. 16.00 200. 238. 38.	7.00 8.00 9.00 10.00 11.00 12.00 13.00 14.00	-80. 60. 230. 390. 510. 560. 450. 330.	-126. 31. 188. 346. 481. 528. 473. 387.	 46. 29. 42. 44. 29. 32. 23. 57.
	15.00 16.00	250. 200.	307. 238.	57. 38.

OBJECTIVE FUNCTION IS 36863.

MODEL PARAMETERS

DEPTH TO TOP OF DYKE IS	1.82
HALF THE THICKNESS IS	3.64
THE ORIGIN IS AT	7.99
THETA-PHI IS	270.00
CONST IS	399.10
REGIONAL AT 1ST STATION IS	184.45
REGIONAL GRADIENT IS	-19.40

********** AT THE END OF 9 ITERATIONS *******

DIS- TANCE	OBSERVED ANOMALY	CALCULATED ANOMALY	ERROR
.00 1.00 2.00 3.00 4.00 5.00 6.00 7.00 8.00 9.00 10.00 11.00 12.00 13.00 14.00 15.00	$\begin{array}{c} -220. \\ -300. \\ -370. \\ -460. \\ -495. \\ -340. \\ -210. \\ -80. \\ 60. \\ 230. \\ 390. \\ 510. \\ 560. \\ 450. \\ 330. \\ 250. \end{array}$	$\begin{array}{c} -217.\\ -287.\\ -375.\\ -472.\\ -494.\\ -366.\\ -204.\\ -57.\\ 82.\\ 222.\\ 372.\\ 515.\\ 543.\\ 445.\\ 342.\\ 260. \end{array}$	3. 13. 5. 12. 1. 26. 6. 23. 22. 8. 18. 5. 17. 5. 12. 10.
16.00	200.	195.	5.

ITERATION NUMBER 9

OBJECTIVE FUNCTION IS 3072.

MODEL PARAMETERS

DEPTH TO TOP OF DYKE IS	1.33
HALF THE THICKNESS IS	3.86
THE ORIGIN IS AT	7.73
THETA-PHI IS	272.90
CONST IS	332.51
REGIONAL AT 1ST STATION IS	127.43
REGIONAL GRADIENT IS	-16.16

-

STOP END OF PROGRAM

SOFTWARE CODING

```
С
      GIODYKE.FOR
      TO REFINE DYKE MODEL PARAMETERS BY GENERALISED INVERSE
С
С
      TECHNIQUE FROM MAGNETIC DATA
      IMPLICIT REAL*8 (A-H,O-Z)
      REAL LAMDA, LAMDA1
      PARAMETER (NM=17, M=17, N=7, NF=7)
      DIMENSION X(17), T(17), TCAL(17), ERROR(17),
                                                    1P(17,7),
      PB(17), DELC(7), PA(1500), DIF(17)
       OPEN(1, FILE='DYUCOMP.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(2, FILE='DYVCOMP.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(3, FILE='DYGENSOL.DAT',
                                       ACCESS='SEQUENTIAL')
                                       ACCESS='SEQUENTIAL')
       OPEN(4, FILE='DYMRES.DAT',
       OPEN(5, FILE='DYDRES.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(6, FILE='DYBCAL.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(7, FILE='DYSVDOUT.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(8, FILE='DYRESYL.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(9, FILE='DYGIO.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(10, FILE='DYKEIN.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(11, FILE='DYKEOUT.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(12, FILE='DYPER.DAT',
                                       ACCESS='SEQUENTIAL')
       OPEN(13, FILE='DYMDIA.DAT',
                                       ACCESS='SEQUENTIAL')
С
       READ(10,801)NOBS, STRIKE, DM, NITER
       READ(10,802) (X(K),K=1,NOBS)
       READ(10,803) (T(K),K=1,NOBS)
       WRITE(8,2)
       WRITE(13,2)
       RAD=0.0174532924
       NN=7
       N2 = 3
       NLAMDA=1
       CON=1.0
       XFAR1=0.0
       XFAR2=0.0
       STR=STRIKE*RAD
       DMR=DM*RAD
       IF(DMR.EO.0.0)THEN
        DM1=1.57079637
      ELSE
        DM1=DATAN(DSIN(STR)/DTAN(DMR))
      ENDIF
С
       LOCATION OF MAXIMUM AND MINIMUM ANOMALIES
       CALL MAXMIN(X,T,NOBS,MAX,MIN,XMAX,XMIN,TMAX,TMIN)
С
       LOCATION OF ORIGIN
      IF(TMIN.NE.0.0) THEN
5
        WIDTH=DABS(XMAX-XMIN)
        D=ORIGIN(X,T,NOBS,MAX,MIN,XMAX,XMIN,TMAX,TMIN)
      ELSE
        FAR=0.5
      WIDTH=FMAX(X,T,NOBS,TMAX,FAR,XFAR1,XFAR2)
```

	D=XMAX
	ENDIF
С	CALCULATION OF INITIAL VALUES OF W,Z AND Q
	R=DABS(TMIN/TMAX)
	TF(R=0.05)10.15.15
10	O=0.0
ŦŬ	2 0.0 7-WTMTH/4 472
1 5	
20	IF(R=0.27)20,25,25
20	Q=30.0
	GO TO 40
25	IF(R-0.7)30,30,35
30	Q=60.0
	GO TO 40
35	Q=90.0
40	S=DSIN(Q*RAD)
	Z=WIDTH*S/(2.0*DSQRT(1+4.0*S*S))
45	IF(XMAX-XMIN)50,60,60
50	TF(TMAX)55.75.75
55	0=0+180 0
55	$Q = Q^{-1} = 0.0$
60	$EV = 10^{-7}$
00 СГ	100.00
05	Q=180.0-Q
	GO TO 75
70	Q=360.0-Q
75	W=2.0*Z
	Q=Q*RAD
	CONST=1.0
	WRITE(11,911)
	DO 80 I=1,9
80	PA(I)=0.0
	DO 85 I=1,3
85	PB(I)=0.0
C	CALCULATION OF INITIAL VALUES OF THE SIZE FACTOR A AND B
0	DO 95 K=1 NOBS
	XX = X(K) - D
	$\pi C \Lambda I (K) = D V K F (XX W Z C O N S T O)$
	EDDOD(1) - V(K)
	ERROR(1) - A(R)
	ERROR(2)=1.0
	ERROR(3) = TCAL(K)
	ERROR(4) = T(K)
	DO 95 J=1,N2
	DO 90 I=1,N2
	L=(J-1)*N2+I
90	PA(L) = PA(L) + ERROR(J) * ERROR(I)
	PB(J) = PB(J) + ERROR(J) * ERROR(I)
95	CONTINUE
	CALL SIMQ(PA, PB, N2, KS)
	A=PB(1)
	B=PB(2)
	CONST=PB(3)
	FIINCT=0

С	CALCULATION OF ANOMALIES OF THE INITIAL MODEL DO 100 K=1,NOBS
	XX=X(K)-D
	TCAL(K) = DYKE(XX, W, Z, CONST, Q) + A * X(K) + B
	DIF(K) = (T(K) - TCAL(K))
	DIF(K) = DABS(DIF(K))
100	FUNCT=FUNCT+(T(K)-TCAL(K))**2
С	INVERSION SCHEME STARTS
	DO 175 ITER=1,NITER
	ICOUNT=0
	ITER1=ITER-1
	QQ=(Q-DM1)/RAD
	WRITE(11,908)ITER1
	WRITE(11,902)
	WRITE(11,903)(X(K),T(K),TCAL(K),DIF(K),K=1,NOBS)
	WRITE(11,906)
	WRITE(11,900)FUNCT
	WRITE $(11,901)Z,W,D,QQ,CONST,B,A$
	IF(FUNCT-NOBS)105,110,110
105	WRITE(11,912)
	STOP
110	IF(ITER-NITER)115,115,180
115	DO 120 L=1,NOBS
	DO 120 MM=1,NN
120	P(L,MM) = 0.0
	C=DCOS(Q)
	S=DSIN(Q)
	Z2=Z*Z
С	CALCULATION OF DERIVATIVES
	DO 130 K=1,NOBS
	R1=X(K)-D
	R2=R1+W
	R3=R1-W
	R4=Z2+R2*R2
	R5=Z2+R3*R3
	R12=R1*R1
	W2=W*W
	R45=R4*R5
	R6=R12-Z2-W2
	R7=R12+Z2+W2
	R8=R12+Z2-W2
	A1=DATAN(R2/Z)
	A2=DATAN(R3/Z)
	ERROR(1)=2.0*CONST*(R7*Z*C-R1*S*R8)/R45
	ERROR(2)=2.0*CONST*W*(R6*C+2.0*Z*R1*S)/R45
	ERROR(3)=2.0*CONST*W*(-R6*S+2.0*Z*R1*C)/R45
	ERROR(4) = -S*(A1-A2) - 0.5*C*DLOG(R4/R5)
	ERROR(4) = ERROR(4) * CONST
	ERROR(5) = (TCAL(K) - A * X(K) - B) / CONST
	ERROR(6) = X(K)
	ERROR(7) = 1.0
	ERROR(8) = T(K) - TCAL(K)

```
PB(K) = ERROR(8)
          P(K, 1) = ERROR(1)
          P(K, 1) = ERROR(1)
          P(K, 2) = ERROR(2)
          P(K,3) = ERROR(3)
          P(K, 4) = ERROR(4)
          P(K, 5) = ERROR(5)
          P(K, 6) = ERROR(6)
          P(K,7) = 1.0
С
     CONSTRUCTION AND SOLUTION OF NORMAL EQUATIONS
130
        CONTINUE
135
        CONTINUE
        ICOUNT=ICOUNT+1
     DO 155 L=1, NOBS
        DO 150 MM=1, NN
        IF( L.EQ.MM) P(L,MM)=P(L,MM)*CON
 150
        CONTINUE
 155
        CONTINUE
         CALL GIO( P, PB, DELC, ITER, NITER, ICOUNT)
      IF(KS.EQ.1) THEN
          WRITE(11,904)
          STOP
         ENDIF
С
      IMPROVEMENT OF THE INITIAL PARAMETERS OF THE MODEL
         WW=0.25*W
      IF(DABS(DELC(1)).GT.WW) DELC(1)=WW*DELC(1)/DABS(DELC(1))
          WD=W+DELC(1)
          ZZ=0.25*Z
      IF(Z.NE.0.0) THEN
        IF(DABS(DELC(2)).GT.ZZ)
           DELC(2) = ZZ*DELC(2)/DABS(DELC(2))
        ENDIF
        ZD=Z+DELC(2)
        DO 160 ID=1,NOBS
          IF(D.GT.X(ID)) THEN
            DX=X(ID+1)-X(ID)
            GO TO 165
          ENDIF
160
        CONTINUE
165
        DDX=0.25*DX
        IF(DABS(DELC(3)).GT.DDX) DELC(3)=DDX*DELC(3)/DABS(DELC(3))
        DD=D+DELC(3)
        IF(DABS(DELC(4)).GT.0.2618)
            DELC(4) = 0.2618 * DELC(4) / DABS(DELC(4))
        OD=O+DELC(4)
        CC=0.50*CONST
        IF(DABS(DELC(5)).GT.CC) DELC(5)=CC*DELC(5)/DABS(DELC(5))
        CD=CONST+DELC(5)
        AD = A + DELC(6)
        BD=B+DELC(7)
        IF(ZD.LT.0.0)ZD=0.001
```

```
FUNCT1=0.0
      CALCULATION OF THE ANOMALIES OF THE IMPROVED MODEL AND
С
С
      OPTIMIZATION
        DO 170 K=1,NOBS
          XX=X(K)-DD
          TCAL(K) = DYKE(XX, WD, ZD, CD, QD) + AD * X(K) + BD
        FUNCT1 = FUNCT1 + (T(K) - TCAL(K)) * *2
        DIF(K) = (T(K) - TCAL(K))
 170
        DIF(K) = DABS(DIF(K))
        IF (FUNCT1.LE.FUNCT) THEN
        Z = ZD
        W=WD
        O=OD
        D=DD
        CONST=CD
        A=AD
        B=BD
        FUNCT=FUNCT1
        NLAMDA=NLAMDA-1
        IF(NLAMDA.EQ.0.0)NLAMDA=1
          LAMDA=0.5*(2**(NLAMDA-1)-1)
          CON=1+LAMDA
        ELSE
        IF(LAMDA.GT.15.0)THEN
          WRITE(11,909)
          STOP
        ENDIF
        LAMDA1=LAMDA
        NLAMDA=NLAMDA+1
        LAMDA=0.5*(2**(NLAMDA-1)-1)
        CON=1+LAMDA
        WRITE(11,910)ITER, FUNCT1, LAMDA1, LAMDA
        GO TO 135
        ENDIF
175
     CONTINUE
180 WRITE(11,905)
     WRITE(8,2)
     WRITE(13,4)
 2
     FORMAT(5X, /'DIAGONAL ELEMENTS OF DATA RESOLUTION ', / )
     FORMAT(5X, /'DIAGONAL ELEMENTS OF MODEL RESOLUTION ',/ )
 4
801
     FORMAT(15,2F10.2,15)
802 FORMAT(8F10.2)
802
     FORMAT(8F10.0)
900
     FORMAT(5X, 'OBJECTIVE FUNCTION IS', F10.0, ///5X,
     'MODEL PARAMETERS'/5X,16('-'))
901
     FORMAT(5X, 'DEPTH TO TOP OF DYKE IS', 5X, F10.2/
     *5X, 'HALF THE THICKNESS IS', 7X, F10.2/
     *5X, 'THE ORIGIN IS AT', 12X, F10.2/
     *5X, 'THETA-PHI IS', 16X, F10.2/
     *5X, 'CONST IS', 20X, F10.2/
     *5X, 'REGIONAL AT 1ST STATION IS', 2X, F10.2/
     *5X, 'REGIONAL GRADIENT IS', 8X, F10.2///)
```

```
902 FORMAT(5X,48('-')/5X,'DIS-
     *',13X, 'OBSERVED',10X,'CALCULATED'/5X,
     * 'TANCE', 12X, 'ANOMALY', 11X, 'ANOMALY', /5X, 48('-'))
       FORMAT(F10.2,8X,F10.0,8X,F10.0,2X,F6.0)
903
       FORMAT(5X,'ILL-CONDITIONED MATRIX, THE',
904
   * 'PROGRAM IS TERMINATED')
905
       FORMAT(5X, 'THE PARAMETERS MAY NOT IMPROVE WITH FURTHER'
     *1X, 'ITERATIONS.'/5X, 'THE LAST PRINTED TABLE
     *2 PROVIDES THE BEST', 1X, 'FITTING SOLUTION.'/5X,
     *3'STOP. END OF PROGRAM')
906
     FORMAT(4X,48('-'))
907
     FORMAT(5X, 'THE OBJECTIVE FUNCTION INCREASES WITH CHANGES IN'
     *1X, 'PARAMETERS.', /5X, 'THE INCREMENTS ARE IGNORED. THE LAST'
     *1X, 'PRINTED TABLE PROVIDES THE SOLUTION')
908
     FORMAT(5X,'ITERATION NUMBER', I4)
909
     FORMAT(5X, 'LAMDA IS INCREASED TO UNUSUALLY',
      *'LARGE VALUE.',/5X,
     * 'THE LAST PRINTED TABLE MAY PROVIDE THE BEST'
     *'FITTING SOLUTION.'
     */5X,'IF NOT, CHECK DATA OR MODEL. THE ITERATION',
     *'IS TERMINATED.')
910
     FORMAT(5X, 'ITERATION NUMBER IS', I4, /5X, 23('-')/
     *5X, 'THE OBJECTIVE FUNCTION IS ', F10.0,/
     *5X, 'THE CURRENT VALUE OF LAMDA IS ', F5.1/5X,
     * 'AS THERE IS AN INCREASE IN THE OBJECTIVE FUNCTION
     *'/5X,'LAMDA IS INCREASED TO',F5.1,
     *' AND THE CALCULATION IS REPEATED.'//)
911
     FORMAT(/5X, 'ANALYSIS BY GENERALISED INVERSE APPROACH ',//
     *5X, 'INTERPRETATION OF MAGNETIC ANOMALIES OF DYKES'
     *1X, 'USING', /5X, 'AN ITERATIVE APPROACH. THE FOLLOWING
     *TABLES PROVIDE', / 5X, 'INFORMATION ON ANOMALIES AND'
     *'MODEL PARAMETERS AT',/5X,'THE END OF EACH
     *ITERATION.',/5X,51('-')//)
912
     FORMAT(5X, 'OBJECTIVE FUNCTION IS LESS THAN THE ALLOWABLE'
     *'ERROR.',/5X,'SOLUTION IN THE LAST PRINTED TABLE MAY BE'
     * 'ACCEPTED.')
      STOP
      END
C SUBROUTINE SIMQ SOLVES A SET OF N-LINEAR EQUATIONS.
C THIS IS ADAPTED FROM THE IBM SUBROUTINE PACKAGE.
С
С
  INPUT
С
   ____
С
     A : MATRIX OF COEFFICIENTS STORED COLUMN-WISE
С
          : VECTOR OF CONSTANTS
      В
С
     N : NUMBER OF EQUATIONS AND VARIABLES
С
C OUTPUT
С
С
    B : COLUMN VECTOR OF FINAL SOLUTION VALUES
```

C C	KS : OUTPUT DIGIT = 0 FOR NORMAL SOLUTUION = 1 FOR SINGULAR MATRIX
	SUBROUTINE SIMQ(A,B,N,KS)
	IMPLICIT REAL*8 (A-H,O-Z)
	DIMENSION A(*),B(*)
	TOL=0.0
	KS=0
	JJ = -N
	DO 40 J=1,N
	JY=J+1
	JJ=JJ+N+1
	BIGA=0
	IT=JJ-J
	DO 10 I=J,N
	IJ=IT+I
	<pre>IF(ABS(BIGA)-ABS(A(IJ))) 5,10,10</pre>
5	BIGA=A(IJ)
	IMAX=I
10	CONTINUE
	IF(ABS(BIGA)-TOL) 15,15,20
15	KS=1
	RETURN
20	I1=J+N*(J-2)
	IT=IMAX-J
	DO 25 K=J,N
	I1=I1+N
	I2=I1+IT
	SAVE=A(I1)
	A(11) = A(12)
	A(I2)=SAVE
25	A(II) = A(II) / BIGA
	SAVE=B(IMAX)
	B(IMAX) = B(J)
	B(J)=SAVE/BIGA
	IF(J-N)30, 45, 30
30	IOS=N*(J-1)
	DO 40 IX=JY, N
	IXJ=IOS+IX
	IT=J-IX
	DO 35 JX=JY,N
	IXJX=N*(JX-1)+IX
	JJX=IXJX+IT
35	A(IXJX) = A(IXJX) - (A(IXJ) * A(JJX))
40	B(IX) = B(IX) - (B(J) * A(IXJ))
45	NY=N-1
	IT=N*N
	DO 50 T = 1.NY
	IA=IT-J
	IB=N-J
	IC=N
	DO 50 K=1,J
	B(IB)=B(IB)-A(IA)*B(IC)

IA=IA-N 50 IC=IC-1 RETURN END

```
C FUNCTION SUBPROGRAM FMAX CALCULATES THE DISTANCE BETWEEN
C TWO POINTS OF EQUAL ANOMALY, REDUCED FROM THEIR MAXIMUM
C VALUE BY A FACTOR
С
С
      INPUT
С
      ____
С
       N : NUMBER OF ANOMALIES IN THE PROFILE
С
       X : DISTANCE OF ANOMALY FROM AN ARBITRARY REFERENCE
       G : ANOMALY VALUES
С
С
   GMAX : AMPLITUDE OF THE MAXIMUM ANOMALY
С
   FAR : FACTOR BY WHICH THE ANOMALY SHOULD BE REDUCED
С
      FUNCTION FMAX(X,G,N,GMAX,FAR,XFAR1,XFAR2)
      IMPLICIT REAL*8 (A-H,O-Z)
      DIMENSION X(*),G(*)
      GH=FAR*GMAX
      K=0
      IF(GMAX)35,5,5
5
     K=K+1
      IF(GH-G(K))15,10,5
10
      XFAR1=X(K)
      GO TO 20
15
      XFAR1=X(K-1)+(X(K)-X(K-1))*(GH-G(K-1))/(G(K)-G(K-1))
20
      K=K+1
      IF(GH-G(K))20,25,30
25
      XFAR2=X(K)
      GO TO 65
30
      XFAR2=X(K-1)+(X(K)-X(K-1))*(GH-G(K-1))/(G(K)-G(K-1))
      GO TO 65
35
     K=K+1
      IF(GH-G(K))35,40,45
40
      XFAR1=X(K)
      GO TO 50
45
      XFAR1=X(K-1)+(X(K)-X(K-1))*(GH-G(K-1))/(G(K)-G(K-1))
50
      K=K+1
      IF(GH-G(K))60,55,50
55
      XFAR2=X(K)
      GO TO 65
60
      XFAR2=X(K-1)+(X(K)-X(K-1))*(GH-G(K-1))/(G(K)-G(K-1))
65
      FMAX=DABS(XFAR2-XFAR1)
      RETURN
      END
```

```
SUBROUTINE MAXMIN CALCULATES THE POSITION AND AMPLITUDES
С
С
   OF MAXIMUN AND MINIMUM ANOMALIES ON A MAGNETIC PROFILE
С
     INPUT
С
       N : NUMBER OF ANOMALIES IN THE PROFILE
С
       X : DISTANCE TO ANOMALY POINTS MEASUR
С
           FROM AN ARBITRARY REFERENCE (ANY UNITS)
С
       T : GEOPHYSICAL ANOMALY
С
     OUTPUT
С
        MAX
                  : STATION NUMBER CLOSEST TO XMAX
С
        MIN
                  : STATION NUMBER CLOSEST TO XMIN
С
                  : POSITION OF THE MAXIMUM ANOMALY
        XMAX
С
                  : POSITION OF MINIMUM ANOMALY
        XMIN
С
                  : AMPLITUDE OF MAXIMUM ANOMALY
        тмах
С
        TMIN
                  : AMPLITUDE OF MINIMUM ANOMALY
С
    SUPPORTING SUBPROGRAM: RINT
      SUBROUTINE MAXMIN(X,T,N,MAX,MIN,XMAX,XMIN,TMAX,TMIN)
      IMPLICIT REAL*8(A-H,O-Z)
      DIMENSION X(*),T(*)
С
      TO FIND THE MAXIMUM ANOMALY
      TMAX=0.0
      DO 10 K=1,N
        IF(DABS(TMAX)-DABS(T(K)))5,5,10
5
        TMAX=T(K)
        MAX=K
10
      CONTINUE
      GR1=T(MAX)-T(MAX-1)
      GR2=T(MAX+1)-T(MAX)
      IF((GR1+GR2).NE.0.0)THEN
        XMAX=0.5*(X(MAX)+X(MAX-1)-GR1*(X(MAX+1)-X(MAX-1)))
        /(GR2-GR1))
        TMAX=RINT(X,N,T,XMAX)
      ELSE
        TMAX=T(MAX)
        XMAX=X(MAX)
      ENDIF
С
      TO FIND THE MINIMUM ANOMALY
      TMIN=0.0
      XMIN=X(N)
      IF(TMAX)15,15,30
15
      DO 25 K=1,N
        IF(TMIN-T(K))20,20,25
20
        TMIN=T(K)
        XMIN=X(K)
        MIN=K
25
      CONTINUE
      GO TO 45
30
      DO 40 K=1,N
        IF(TMIN-T(K))40,35,35
35
          TMIN=T(K)
          XMIN=X(K)
          MIN=K
40
      CONTINUE
```

```
45
      IF(TMIN)50,100,50
50
      GR1=T(MIN)-T(MIN-1)
      GR2=T(MIN+1)-T(MIN)
      IF((GR1+GR2).NE.0.0)THEN
        XMIN=0.5*(X(MIN)+X(MIN-1)-GR1*(X(MIN+1)-X(MIN-1))
       /(GR2-GR1))
      TMIN=RINT(X,N,T,XMIN)
      ELSE
        TMIN=T(MIN)
        XMIN=X(MIN)
      ENDIF
100
     RETURN
      END
C FUNCTION SUBPROGRAM DYKE CALCULATES THE MAGNETIC ANOMALY
C (IN GAMMAS) OF AN ARBITRARILY MAGNETIZED DYKE
С
     INPUT
С
        X : DISTANCE OF POINT OF CALCULATION
             FROM THE CENTRE OF THE DYKE (ANY UNITS)
С
С
          : HALF WIDTH OF THE DYKE (SAME UNITS AS X)
        W
С
        Z : DEPTH TO THE DYKE (SAME UNITS AS X )
С
      SIZE : SIZE FACTOR, BEING A FUNCTION OF
С
             INTENSITY OF MAGNETIZATION, (IN GAMMAS)
С
             AND DIP OF THE DYKE
С
          : ANGLE, BEING A FUNCTION OF DIRECTION
        0
С
             OF MAGNETIZATION, DIP OF THE DYKE AND
С
             THE DIRECTION OF MEASUREMENT (IN RADIANS)
С
      FUNCTION DYKE(X,W,Z,SIZE,O)
      IMPLICIT REAL*8 (A-H,O-Z)
      PIBY2=1.57079632
      R=1.0E-03
      C=DCOS(Q)
      S=DSIN(Q)
      Z2=Z*Z
      R1=X+W
      R2=X-W
      R3=Z2+R1*R1
      R4=Z2+R2*R2
      ABSX=DABS(X)
      IF(Z.NE.0.0)THEN
        ANG=DATAN(R1/Z) - DATAN(R2/Z)
      ELSEIF(ABSX.EQ.W)THEN
        ANG=PIBY2
      ELSE
        ANG=PIBY2*(1-(ABSX-W)/DABS(ABSX-W))
      ENDIF
      IF(R3.LT.R)R3=R
      IF(R4.LT.R)R4=R
```

RTERM=DLOG(R3/R4) DYKE=SIZE*(C*ANG-0.5*S*RTERM) RETURN END FUNCTION SUBPROGRAM ORIGIN LOCATES THE POSITION OF THE С С DYKE'S CENTRE INPUT: С С N : NUMBER OF ANOMALIES IN THE PROFILE С X : DISTANCE OF ANOMALY FROM AN ARBITRARY REFERENCE С T : MAGNETIC ANOMALY (GAMMAS) С XMAX : POSITION OF MAXIMUM ANOMALY С XMIN : POSITION OF MINIMUM ANOMALY С TMAX : AMPLITUDE OF MAXIMUM ANOMALY С TMIN : AMPLITUDE OF MINIMUM ANOMALY С MAX : STATION NUMBER CLOSEST TO XMAX С MIN : STATION NUMBER CLOSEST TO XMIN FUNCTION ORIGIN(X,T,N,MAX,MIN,XMAX,XMIN,TMAX,TMIN) IMPLICIT REAL*8 (A-H,O-Z) DIMENSION X(*),T(*) T0=TMAX+TMIN IF(XMAX-XMIN)50,5,5 5 IF(TMAX)10,30,30 10 K=MIN+1 15 K=K+1IF(T0-T(K))15,20,25 20 ORIGIN=X(K) GO TO 95 25 ORIGIN=X(K-1)+(X(K)-X(K-1))*(T0-T(K-1))/(T(K)-T(K-1))GO TO 95 30 K=MIN+1 K=K+135 IF(T0-T(K))45,40,35 40 ORIGIN=X(K) GO TO 95 45 ORIGIN=X(K-1)+(X(K)-X(K-1))*(T0-T(K-1))/(T(K)-T(K-1))GO TO 95 50 IF(TMAX)55,75,75 55 K=MAX+1 60 K=K+1IF(T0-T(K))70,65,60 65 ORIGIN=X(K) GO TO 95 70 ORIGIN=X(K-1)+(X(K)-X(K-1))*(T0-T(K-1))/(T(K)-T(K-1))GO TO 95 75 K=MAX+1 80 K=K+1IF(T0-T(K))80,85,90 85 ORIGIN=X(K) GO TO 95 ORIGIN=X(K-1)+(X(K)-X(K-1))*(T0-T(K-1))/(T(K)-T(K-1))90

95 CONTINUE RETURN END

C FUNCTION SUBPROGRAM RINT CALCULATES BY INTERPOLATION C THE VALUE OF A FUNCTION F(X) FROM TABULATED FUNCTION C VALUES AT AN INTERMEDIATE VALUE OF X C C INPUT : C -----C N : NUMBER OF DIGITIZED FUNCTION VALUES C X : ARRAY OF VALUES OF X AGAINST WHICH X

С IS NOTED С FX : ARRAY OF VALUES OF F(X)С XX : VALUE OF X AT WHICH F(X) NEEDS TO BE С CALCULATED FUNCTION RINT(X,N,FX,XX) IMPLICIT REAL*8(A-H,O-Z) DIMENSION X(*), FX(*) DO 15 K=1,N IF(XX-X(K))5,10,15 5 M=K-1 GO TO 20 10 RINT=FX(K) RETURN 15 CONTINUE 20 IF((XX-X(M)).GT.(X(M+1)-XX))M=M+1IF(M.LE.2.OR.M.GT.(N-2))THEN RINT=FX(M)RETURN ENDIF X1=X(M-2)X2 = X(M - 1)X0 = X(M)X3=X(M+1)X4 = X(M+2)T1 = FX(M-2)T2=FX(M-1)TO = FX(M)T3=FX(M+1)T4=FX(M+2)A = (XX - X1) * (XX - X2) * (XX - X0) * (XX - X4) * (XX - X3)B=T1/((XX-X1)*(X1-X2)*(X1-X3)*(X1-X4)*(X1-X0)) C=T2/((XX-X2)*(X2-X1)*(X2-X3)*(X2-X4)*(X2-X0)) D=T0/((XX-X0)*(X0-X1)*(X0-X2)*(X0-X3)*(X0-X4))E=T3/((XX-X3)*(X3-X1)*(X3-X2)*(X3-X4)*(X3-X0))F=T4/((XX-X4)*(X4-X1)*(X4-X2)*(X4-X3)*(X4-X0)) RINT=A*(B+C+D+E+F)RETURN END

SUBROUTINE GIO(A, D, DELC, ITR, NITR, ICOUNT) С IMPLICIT DOUBLE PRECISION (A-H,O-Z) IMPLICIT REAL*8(A-H,O-Z) PARAMETER (NM=26, M=26, N=7, NF=7)DIMENSION A(NM,NM),W(NM),U(NM,NM),V(NM,NM),RV1(NM) DIMENSION WW(NM), UU(NM, NM), VV(NM, NM), CT(NM, NM), DELC(NM) DIMENSION (NM,NM),C(NM,NM),UT(NM,NM),VT(NM,NM),ILT(NM),D(NM) LOGICAL MATU, MATV VAR=1.0 MATU=.TRUE. MATV=.TRUE. IERR=0 CALL SVD(NM, M, N, A, WW, MATU, UU, MATV, VV, IERR, RV1) С WRITE(2,*)'Unordered - decomposition ' С CALL OUTPUT(A, UU, VV, WW, NM, M, N, NF) DO 5 I=1,N W(I) = -WW(I)5 ILT(I) = ICALL ARRANGE(W, ILT, NM, N) DO 6 I=1,N 6 W(I) = -W(I)DO 7 J=1,N K = ILT(J)DO 7 I=1,M 7 U(I,J) = UU(I,K)DO 8 J=1,N K = ILT(J)DO 8 I=1,N 8 V(I,J) = VV(I,K)C*** WRITE(7,*)'ordered - decomposition ' CALL OUTPUT(A,U,V,W,NM,M,N,NF,ITR,NITR) С Computing the closeness ratio SUM=0.0DO 14 I=1,N 14 SUM=SUM+W(I) SUM1=0.0 DO 100 II=1,NF SUM1=SUM1+W(II) CR=SUM1/SUM*100.0 IF(ITR.NE.NITR) GO TO 9999 WRITE(3,2)II, CR, ITR, ICOUNT WRITE(4,2)II, CR, ITR, ICOUNT WRITE(5,2)II, CR, ITR, ICOUNT WRITE(6,2)II, CR, ITR, ICOUNT WRITE(7,2)II, CR, ITR, ICOUNT WRITE(9,2)II, CR, ITR, ICOUNT WRITE(12,22)II, CR, ITR, ICOUNT 22 FORMAT(1X, I3, F8.4, 1X, I3, 1X, I3) 2 FORMAT(/1X, 'No.factrs cnsidered = ',I3/1X,

^{1&#}x27;Closeness ratio =',F10.3,3X,'ITR:', I3,3X,'ICOUNT:',I3/)

```
9999 CONTINUE
      DO 1 I=1,M
      DO 1 J=1,II
1
      X(I,J)=U(I,J)*W(J)
      CALL TRNSPZ(V,VT,NM,N,II)
      CALL MATMLT(X,VT,C,NM,M,N,II)
      IF(ITR.NE.NITR) GO TO 8888
      WRITE(6,*)'Back calculatin matrix '
      DO 332 I=1, M
      DO 332 j=1, N
 332
     WRITE(6,333)I,J,C(I,J)
 333
     FORMAT(1X, I3, 1X, I3, 1X, E12.5)
8888
      CONTINUE
      CALL TRNSPZ(U,UT,NM,M,II)
      CALL MATMLT(U,UT,C,NM,M,M,II)
      IF(ITR.NE.NITR) GO TO 7777
      WRITE(5,*)'Data - Resolution matrix '
      DO 331 I=1, M
      DO 331 J=1, M
 331 WRITE(5,333)I,J,C(I,J)
      WRITE(8,222)( C(L,L), L=1, M )
 222
     FORMAT(1X,20F6.3)
7777
      CONTINUE
      CALL MATMLT(V,VT,C,NM,N,N,II)
      IF(ITR.NE.NITR) GO TO 6666
      WRITE(4,*)'Model - Resolution matrix '
      DO 330 I=1, N
      DO 330 J=1, N
 330 WRITE(4,333)I,J, C(I,J)
6666
      CONTINUE
С
      Constructing of the generalized inverse operator
С
       WRITE(7,*)'Generalized inverse operator '
      DO 9 I=1,N
      DO 9 J=1,II
9
      X(I,J) = V(I,J) / W(J)
      CALL MATMLT(X,UT,C,NM,N,M,II)
      IF(ITR.NE.NITR) GO TO 5555
      WRITE(9,*) ' GENERALISED INVERSE OPERATOR '
      DO 335 I=1, N
      DO 335 J=1, M
     WRITE(9,333) I,J,C(I,J)
 335
5555
      CONTINUE
С
      Computing the generalized inverse solution
С
       WRITE(7,*)'Generalized inverse solution '
      WRITE(7,*)'Generalized inverse solution '
      DO 11 I=1,N
      DELC(I)=0.0
      DO 11 J=1,M
11
      DELC(I) = DELC(I) + C(I,J) * D(J)
      WRITE(7, 12)(DELC(I), I=1, N)
12
      FORMAT(7(E10.4,1X))
      Computing the covariance of the solutioon
С
```

С	WRITE(7,*)'covariance of the solution '
	CALL TRNSPZ(C,CT,NM,N,M)
	DO 13 I=1,M
	DO 13 J=1,N
13	X(I,J) = C(I,J) * VAR
	CALL MATMLT(X,CT,C,NM,N,N,M)
100	CONTINUE
	RETURN
	END
SUBRO	UTINE SVD(NM,M,N,A,W,MATU,U,MATV,V,IERR,RV1)
	IMPLICIT REAL*8(A-H,O-Z)
С	IMPLICIT DOUBLE PRECISION (A-H,O-Z)

- C IMPLICIT DOUBLE PRECISION (A-H, O-Z) DIMENSION A(NM,NM),W(NM),U(NM,NM),V(NM,NM),RV1(NM) LOGICAL MATU,MATV C THIS SUBROUTINE IS A TRANSLATION OF THE ALGOL PROCEDURE SVD,
- C NUM. MATH. 14, 403-420(1970) BY GOLUB AND REINSCH.
- C HANDBOOK FOR AUTO. COMP., VOL II-LINEAR ALGEBRA, 134-151(1971).
- C THIS SUBROUTINE DETERMINES THE SINGULAR VALUE DECOMPOSITION C T
- C A=USV OF A REAL M BY N RECTANGULAR MATRIX. HOUSEHOLDER
 C BIDIAGONALIZATION AND A VARIANT OF THE QR ALGORITHM ARE USED.
 C
 - ON INPUT.

С

С

С

С

- C NM MUST BE SET TO THE ROW DIMENSION OF TWO-DIMENSIONAL
 C ARRAY PARAMETERS AS DECLARED IN THE CALLING PROGRAM
 C DIMENSION STATEMENT. NOTE THAT NM MUST BE AT LEAST
 C AS LARGE AS THE MAXIMUM OF M AND N.
- C M IS THE NUMBER OF ROWS OF A (AND U).
- C N IS THE NUMBER OF COLUMNS OF A (AND U) AND THE ORDER OF V. C A CONTAINS THE RECTANGULAR INPUT MATRIX TO BE DECOMPOSED.
- MATU SHOULD BE SET TO .TRUE. IF THE U MATRIX IN THE
 DECOMPOSITION IS DESIRED, AND TO .FALSE. OTHERWISE.
 MATV SHOULD BE SET TO .TRUE. IF THE V MATRIX IN THE
 DECOMPOSITION IS DESIRED, AND TO .FALSE. OTHERWISE.
 ON OUTPUT.
- C A IS UNALTERED (UNLESS OVERWRITTEN BY U OR V).
 C W CONTAINS THE N (NON-NEGATIVE) SINGULAR VALUES OF A (THE
 C DIAGONAL ELEMENTS OF S). THEY ARE UNORDERED. IF AN
 C ERROR EXIT IS MADE, THE SINGULAR VALUES SHOULD BE CORRECT
 C FOR INDICES IERR+1, IERR+2,...,N.
 C U CONTAINS THE MATRIX U (ORTHOGONAL COLUMN VECTORS) OF THE
- C DECOMPOSITION IF MATU HAS BEEN SET TO .TRUE. OTHERWISE C U IS USED AS A TEMPORARY ARRAY. U MAY COINCIDE WITH A.
- C IF AN ERROR EXIT IS MADE, THE COLUMNS OF U CORRESPONDING

```
С
      TO INDICES OF CORRECT SINGULAR VALUES SHOULD BE CORRECT.
С
С
      V CONTAINS THE MATRIX V (ORTHOGONAL) OF THE DECOMPOSITION IF
С
      MATV HAS BEEN SET TO .TRUE. OTHERWISE V IS NOT REFERENCED.
С
      V MAY ALSO COINCIDE WITH A IF U IS NOT NEEDED.
                                                       IF AN ERROR
      EXIT IS MADE, THE COLUMNS OF V CORRESPONDING TO INDICES OF
С
С
      CORRECT SINGULAR VALUES SHOULD BE CORRECT.
С
      IERR IS SET TO
С
      ZERO
                 FOR NORMAL RETURN,
С
      Κ
                 IF THE K-TH SINGULAR VALUE HAS NOT BEEN
С
                 DETERMINED AFTER 30 ITERATIONS.
С
С
      RV1 IS A TEMPORARY STORAGE ARRAY.
С
С
      THIS IS A MODIFIED VERSION OF A ROUTINE FROM THE EISPACK
С
      COLLECTION BY THE NATS PROJECT
С
С
      MODIFIED TO ELIMINATE MACHEP
      IERR = 0
      DO 100 I = 1, M
С
      DO 100 J = 1, N
            U(I,J) = A(I,J)
100
      CONTINUE
С
      ..... HOUSEHOLDER REDUCTION TO BIDIAGONAL FORM
      . . . . . . . . . .
      G = 0.0D0
      SCALE = 0.0D0
      ANORM = 0.0D0
С
      DO 300 I = 1, N
         L = I + 1
         RV1(I) = SCALE * G
         G = 0.0D0
         S = 0.0D0
         SCALE = 0.0D0
         IF (I .GT. M) GO TO 210
С
         DO 120 K = I, M
  120
         SCALE = SCALE + DABS(U(K,I))
С
         IF (SCALE .EQ. 0.0D0) GO TO 210
С
         DO 130 K = I, M
            U(K,I) = U(K,I) / SCALE
            S = S + U(K,I) * * 2
  130
         CONTINUE
С
         F = U(I,I)
         G = -DSIGN(DSQRT(S), F)
         H = F * G - S
         U(I,I) = F - G
```

```
IF (I .EQ. N) GO TO 190
С
         DO 150 J = L, N
            S = 0.0D0
С
            DO 140 K = I, M
  140
            S = S + U(K,I) * U(K,J)
С
            F = S / H
С
            DO 150 K = I, M
                U(K,J) = U(K,J) + F * U(K,I)
  150
         CONTINUE
  190
         DO 200 K = I, M
  200
         U(K,I) = SCALE * U(K,I)
С
  210
         W(I) = SCALE * G
         G = 0.0D0
         S = 0.0D0
         SCALE = 0.0D0
         IF (I .GT. M .OR. I .EQ. N) GO TO 290
         DO 220 K = L, N
  220
         SCALE = SCALE + DABS(U(I,K))
С
         IF (SCALE .EQ. 0.0D0) GO TO 290
         DO 230 K = L, N
            U(I,K) = U(I,K) / SCALE
            S = S + U(I,K) * 2
  230
         CONTINUE
С
         F = U(I,L)
         G = -DSIGN(DSQRT(S), F)
         H = F * G - S
         U(I,L) = F - G
С
         DO 240 K = L, N
  240
         RV1(K) = U(I,K) / H
С
         IF (I .EQ. M) GO TO 270
С
         DO 260 J = L, M
            S = 0.0D0
С
            DO 250 K = L, N
  250
            S = S + U(J,K) * U(I,K)
С
            DO 260 \text{ K} = \text{L}, \text{N}
               U(J,K) = U(J,K) + S * RV1(K)
  260
         CONTINUE
С
         DO 280 K = L, N
  270
  280
         U(I,K) = SCALE * U(I,K)
```

С

```
290
         ANORM = DMAX1(ANORM, DABS(W(I))+DABS(RV1(I)))
  300
         CONTINUE
С
      ..... ACCUMULATION OF RIGHT-HAND TRANSFORMATIONS
      . . . . . . . . . .
      IF (.NOT. MATV) GO TO 410
      ..... FOR I=N STEP -1 UNTIL 1 DO -- .....
С
      DO 400 \text{ II} = 1, \text{ N}
         I = N + 1 - II
         IF (I .EQ. N) GO TO 390
         IF (G .EQ. 0.0D0) GO TO 360
С
         DO 320 J = L, N
      ..... DOUBLE DIVISION AVOIDS POSSIBLE UNDERFLOW
С
      . . . . . . . . . .
  320
         V(J,I) = (U(I,J) / U(I,L)) / G
С
         DO 350 J = L, N
            S = 0.0D0
            DO 340 K = L, N
  340
            S = S + U(I,K) * V(K,J)
С
            DO 350 K = L, N
               V(K,J) = V(K,J) + S * V(K,I)
  350
         CONTINUE
С
  360
         DO 380 J = L, N
            V(I,J) = 0.0D0
            V(J,I) = 0.0D0
  380
         CONTINUE
С
  390
         V(I,I) = 1.0D0
         G = RV1(I)
         L = I
  400 CONTINUE
      ..... ACCUMULATION OF LEFT-HAND TRANSFORMATIONS
С
       . . . . . . . . . .
  410 IF (.NOT. MATU) GO TO 510
С
      .....FOR I=MIN(M,N) STEP -1 UNTIL 1 DO -- .....
      MN = N
      IF (M .LT. N) MN = M
С
      DO 500 II = 1, MN
         I = MN + 1 - II
         L = I + 1
         G = W(I)
         IF (I .EQ. N) GO TO 430
С
         DO 420 J = L, N
  420
         U(I,J) = 0.0D0
С
  430
         IF (G .EQ. 0.0D0) GO TO 475
```

```
IF (I .EQ. MN) GO TO 460
С
         DO 450 J = L, N
            S = 0.0D0
С
            DO 440 \text{ K} = \text{L}, \text{M}
  440
            S = S + U(K,I) * U(K,J)
С
      ..... DOUBLE DIVISION AVOIDS POSSIBLE UNDERFLOW
         . . . . . . . .
            F = (S / U(I,I)) / G
С
            DO 450 \text{ K} = \text{I}, \text{M}
               U(K,J) = U(K,J) + F * U(K,I)
  450
         CONTINUE
С
         DO 470 J = I, M
  460
  470
         U(J,I) = U(J,I) / G
С
         GO TO 490
        DO 480 J = I, M
  475
  480
         U(J,I) = 0.0D0
С
  490
        U(I,I) = U(I,I) + 1.0D0
  500 CONTINUE
      ..... DIAGONALIZATION OF THE BIDIAGONAL FORM .....
С
      ..... FOR K=N STEP -1 UNTIL 1 DO -- .....
С
  510 DO 700 KK = 1, N
        K1 = N - KK
         K = K1 + 1
         ITS = 0
     ..... TEST FOR SPLITTING.
С
С
                FOR L=K STEP -1 UNTIL 1 DO -- .....
  520
         DO 530 LL = 1, K
            L1 = K - LL
            L = L1 + 1
            IF (DABS(RV1(L)) + ANORM .EQ. ANORM) GO TO 565
      ..... RV1(1) IS ALWAYS ZERO, SO THERE IS NO EXIT
С
С
                 THROUGH THE BOTTOM OF THE LOOP .....
            IF (DABS(W(L1)) + ANORM .EQ. ANORM) GO TO 540
  530
         CONTINUE
      ..... CANCELLATION OF RV1(L) IF L GREATER THAN 1
С
      . . . . . . . . . .
  540
        C = 0.0D0
         S = 1.0D0
С
         DO 560 I = L, K
            F = S * RV1(I)
            RV1(I) = C * RV1(I)
            IF (DABS(F) + ANORM .EQ. ANORM) GO TO 565
            G = W(I)
            H = DSORT(F*F+G*G)
            W(I) = H
```

C = G / HS = -F / HIF (.NOT. MATU) GO TO 560 С DO 550 J = 1, M Y = U(J, L1)Z = U(J,I)U(J,L1) = Y * C + Z * SU(J,I) = -Y * S + Z * C550 CONTINUE С 560 CONTINUE TEST FOR CONVERGENCE С 565 Z = W(K)IF (L .EQ. K) GO TO 650 SHIFT FROM BOTTOM 2 BY 2 MINOR С IF (ITS .EQ. 30) GO TO 1000 ITS = ITS + 1X = W(L)Y = W(K1)G = RV1(K1)H = RV1(K)F = ((Y - Z) * (Y + Z) + (G - H) * (G + H)) / (2.0D0 * H)* Y) G = DSQRT(F*F+1.0D0)F = ((X - Z) * (X + Z) + H * (Y / (F + DSIGN(G,F)) - H))/XС NEXT QR TRANSFORMATION C = 1.0D0S = 1.0D0С DO 600 I1 = L, K1 I = I1 + 1G = RV1(I)Y = W(I)H = S * GG = C * GZ = DSQRT(F*F+H*H)RV1(I1) = ZC = F / ZS = H / ZF = X * C + G * SG = -X * S + G * CH = Y * SY = Y * CIF (.NOT. MATV) GO TO 575 DO 570 J = 1, N X = V(J, I1)Z = V(J,I)V(J,I1) = X * C + Z * SV(J,I) = -X * S + Z * C570 CONTINUE

-

С

```
575 Z = DSQRT(F*F+H*H)
           W(I1) = Z
    ..... ROTATION CAN BE ARBITRARY IF Z IS ZERO .....
С
           IF (Z .EQ. 0.0D0) GO TO 580
           C = F / Z
           S = H / Z
           F = C * G + S * Y
 580
           X = -S * G + C * Y
           IF (.NOT. MATU) GO TO 600
           DO 590 J = 1, M
              Y = U(J, I1)
              Z = U(J,I)
              U(J,I1) = Y * C + Z * S
              U(J,I) = -Y * S + Z * C
  590
           CONTINUE
С
 600
        CONTINUE
        RV1(L) = 0.0D0
        RV1(K) = F
        W(K) = X
        GO TO 520
      ..... CONVERGENCE .....
С
        IF (Z .GE. 0.0D0) GO TO 700
 650
С
     ..... W(K) IS MADE NON-NEGATIVE .....
        W(K) = -Z
        IF (.NOT. MATV) GO TO 700
        DO 690 J = 1, N
 690
        V(J,K) = -V(J,K)
 700 CONTINUE
     GO TO 1001
      ..... SET ERROR -- NO CONVERGENCE TO A
С
С
                SINGULAR VALUE AFTER 30 ITERATIONS ...
1000 \text{ IERR} = K
1001 RETURN
     END
```